

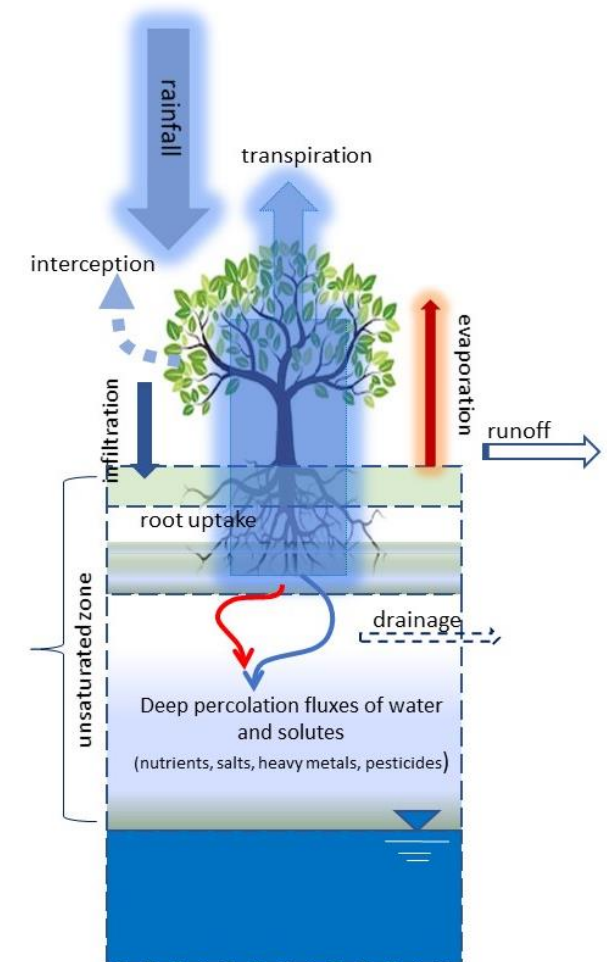
# FLOWS Agrohydrological Model

Theoretical basis



# The model

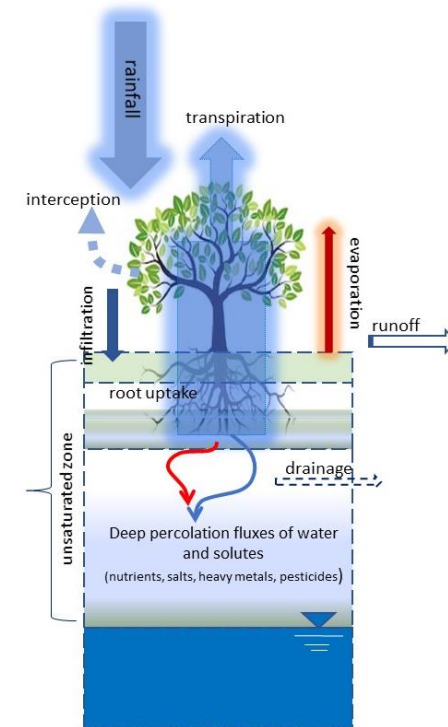
- It is a physically-based model to simulate water flow and solute transport (and transformations) in the soil in the presence of vegetation.
- In the scope of NPP-SOL project: it is the agro-hydrological model to evaluate water and nutrient fluxes under different scenarios.



# The model produces, node by node,

# information on the time evolution of (among many other outputs):

- Soil water contents and pressure potentials in the soil profile;
- Solute (tracers, adsorbed and reactive solutes) concentrations in the soil profile. The model also allows for nitrogen transport simulations by solving the ADE twice, once for N-NH<sub>4</sub> and once for N-NO<sub>3</sub>, with appropriate exchange terms;
- Water and solute uptake and actual evapotranspiration for simulated crops;
- Deep percolation water fluxes
- Deep percolation solute fluxes;
- Root uptake of water and solutes;
- Water and solute fluxes to runoff;
- Drainage water fluxes and related solute fluxes.
- Irrigation fluxes computed by the model;
- Temperature in the soil profile



# Water flow equation

- Richards Equation for 1-D soil water flow:

$$C(h) \frac{\partial h}{\partial t} = \frac{\partial}{\partial z} \left( K(h) \frac{\partial h}{\partial z} - K(h) \right) - S_w$$

where  $C(h)=d\theta/dh$  [ $L^{-1}$ ] is the soil water capacity,  $h$  [L] is the soil water pressure head,  $t$  [T] is time,  $z$  [L] is the vertical coordinate being positive upward,  $K(h)$  [ $L T^{-1}$ ] the hydraulic conductivity and  $S_w$  [ $T^{-1}$ ] is a sink term describing water uptake by plant roots,  $S_r$ , and/or lateral water drainage,  $S_{dr}$ , so that  $S_w=S_r+S_{dr}$ .

# Solute Transport Equation

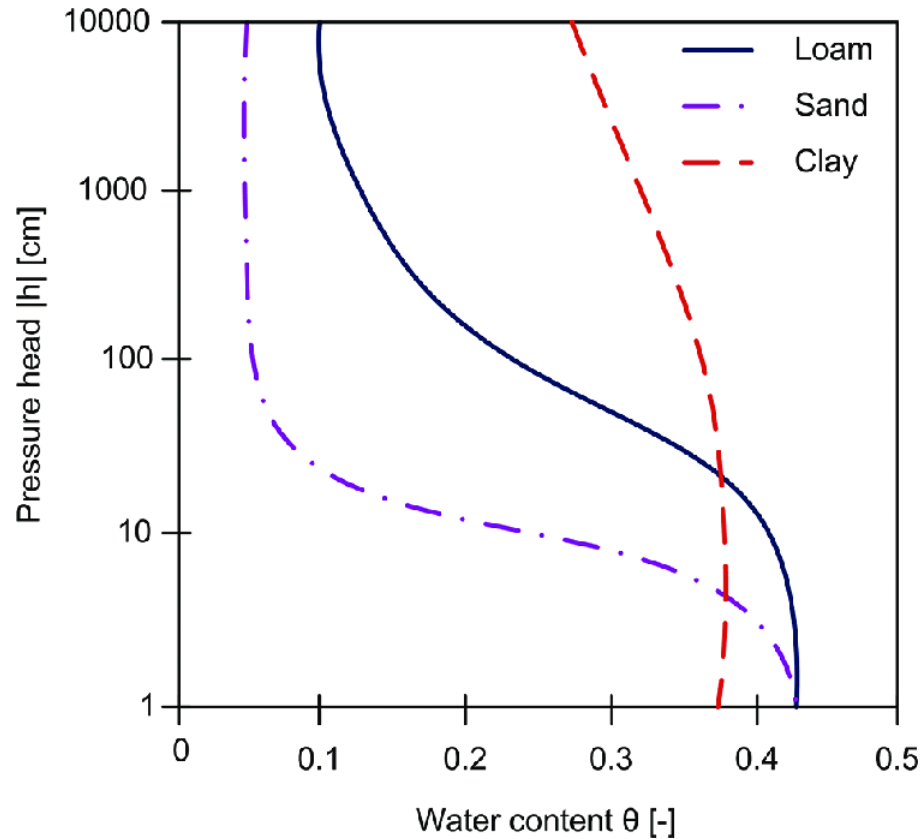
- The 1-D form of the Advection-Dispersion Equation:

$$\frac{\partial \theta C}{\partial t} + \rho_b \frac{\partial C_s}{\partial t} + \frac{\partial \theta_g C_g}{\partial t} = -\frac{\partial q C}{\partial z} + \frac{\partial}{\partial z} \left( \theta D_h \frac{\partial C}{\partial z} \right) + \frac{\partial}{\partial z} \left( \theta_g D_g^s K_H \frac{\partial C}{\partial z} \right) - S_s$$

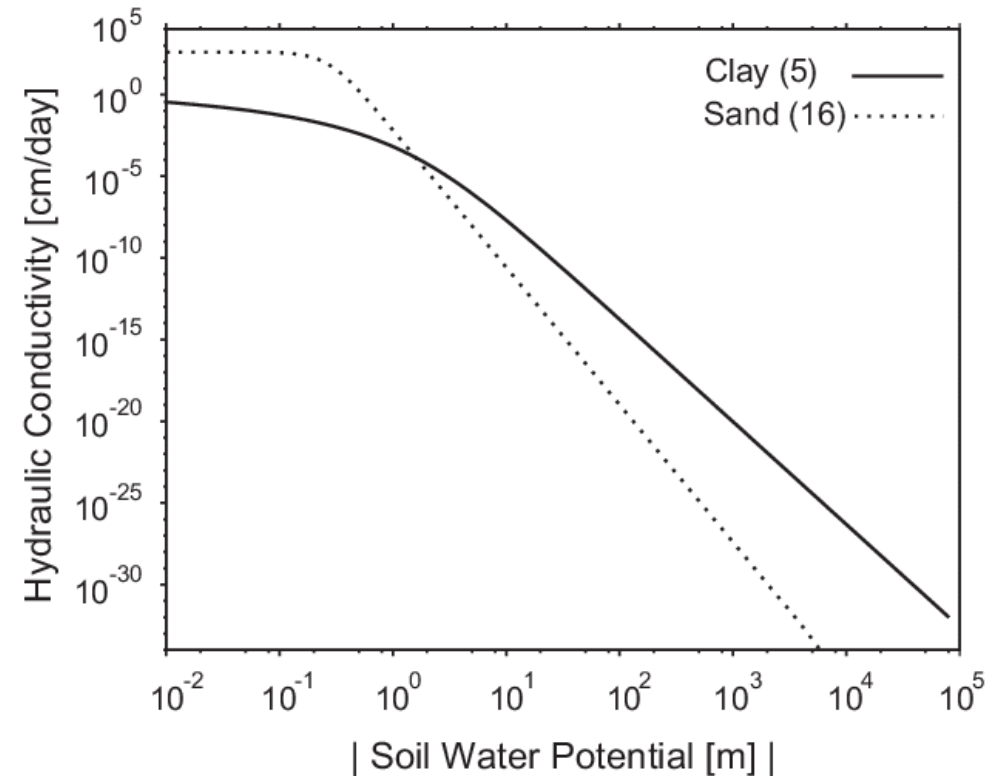
In the equation,  $C$  [ $M L^{-3}$ ],  $C_s$  [ $M M^{-1}$ ] and  $C_g$  [ $M L^{-3}$ ], are the amount of solute in the liquid, adsorbed and gaseous phases, respectively,  $q$  [ $L T^{-1}$ ] is the darcian water flux,  $\rho_b$  [ $M L^{-3}$ ] is the bulk density,  $D_h$  [ $L^2 T^{-1}$ ] the hydrodynamic dispersion coefficient,  $D_g^s$  is the dispersion coefficient in the gaseous phases [ $L^2 T^{-1}$ ],  $\theta_g$  is the volumetric air content in soil,  $S_s$  [ $ML^{-3}T^{-1}$ ] is a source-sink term for solutes,  $K_H$  is the dimensionless Henry constant.

# Richards Equation Requirements: Soil Hydraulic Properties (SHP)

- Water Retention Curve,  $\theta(h)$



- Hydraulic Conductivity Curve,  $K(h)$



## SHP Models Available in FLOWS

FLOWS requires the SHP for each soil horizon. The soil hydraulic parameters that can be used are those of:

- Unimodal van Genuchten – Mualem model;
- Unimodal Gardner & Russo;
- Bimodal Durner – Mualem model; or
- Bimodal Ross & Smettem

## Example: van Genuchten – Mualem

van Genuchten (1981) water retention  $S_e = \frac{\theta - \theta_r}{\theta_s - \theta_r} = \left[ 1 + |\alpha_{VG} h|^n \right]^{-m}$

Mualem (1976) hydraulic conductivity  $K_r(S_e) = \frac{K(S_e)}{K_0} = S_e^\tau \left[ 1 - \left( 1 - S_e^{1/m} \right)^m \right]^2$



# Root Uptake, root distribution and ET

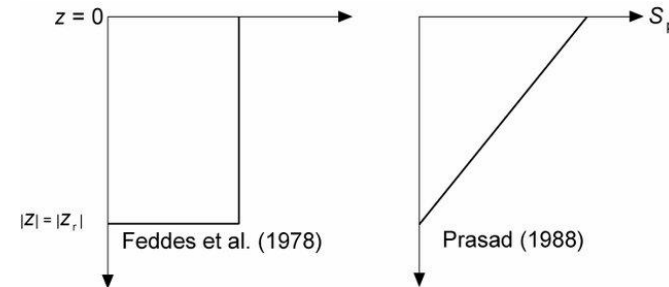
(macroscopic approach)

$$ET_r \text{ \& } K_c \rightarrow ET_p \rightarrow LAI \rightarrow E_p + T_p$$

$$S_{r,p}(z) = g(z)T_p$$

$$T_p = \int_0^{Dr} S_{r,p}(z) dz$$

- $g(z)$  {
1. Feddes (uniform)
  2. Prasad (triangular)
  3. Vrugt
  4. Logistic



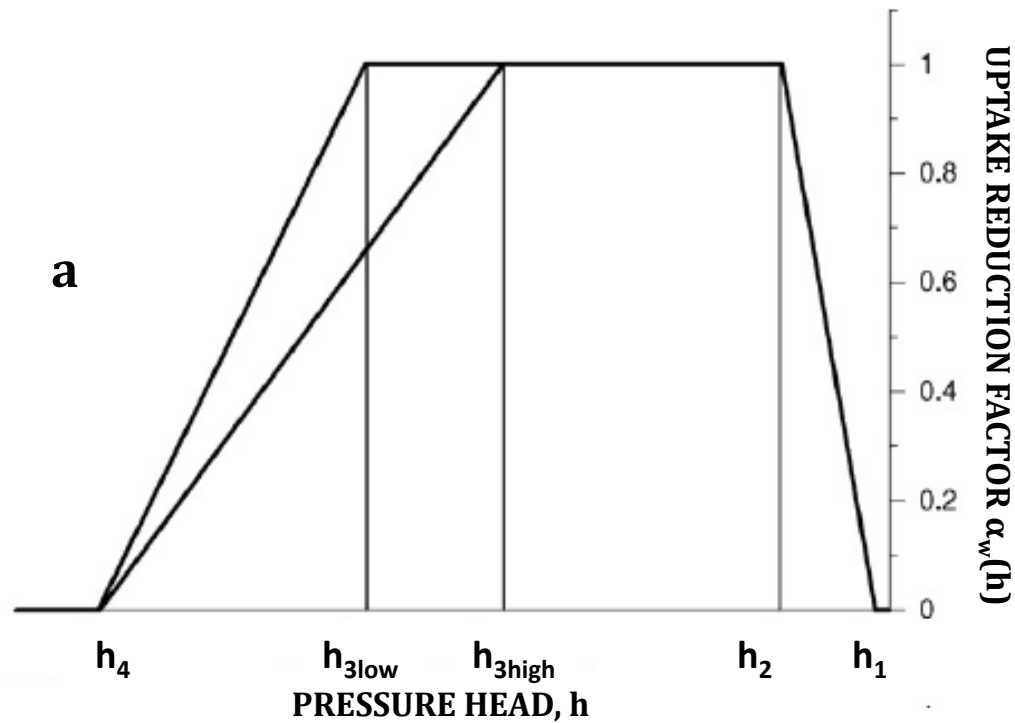
Water Uptake reduction coming from both water and osmotic stresses

$$S_{r,a} = \alpha_w(h)\alpha_s(h_{os})S_p = \alpha_w(h)\alpha_s(h_{os})g(z)T_p$$

# Water stress Uptake reduction function

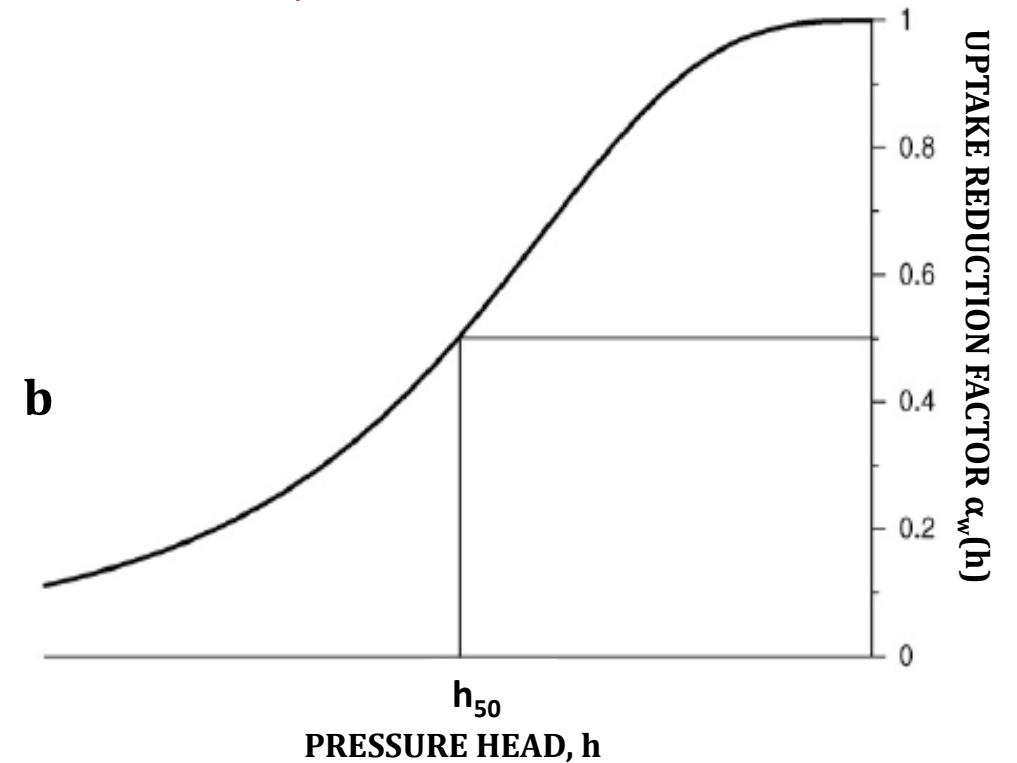
$$\alpha(h) = \begin{cases} \frac{h - h_4}{h_3 - h_4}, & h_3 > h > h_4 \\ 1, & h_2 \geq h \geq h_3 \\ \frac{h - h_1}{h_2 - h_1}, & h_1 > h > h_2 \\ 0, & h \leq h_4 \text{ or } h \geq h_1 \end{cases}$$

Feddes Uptake reduction factor



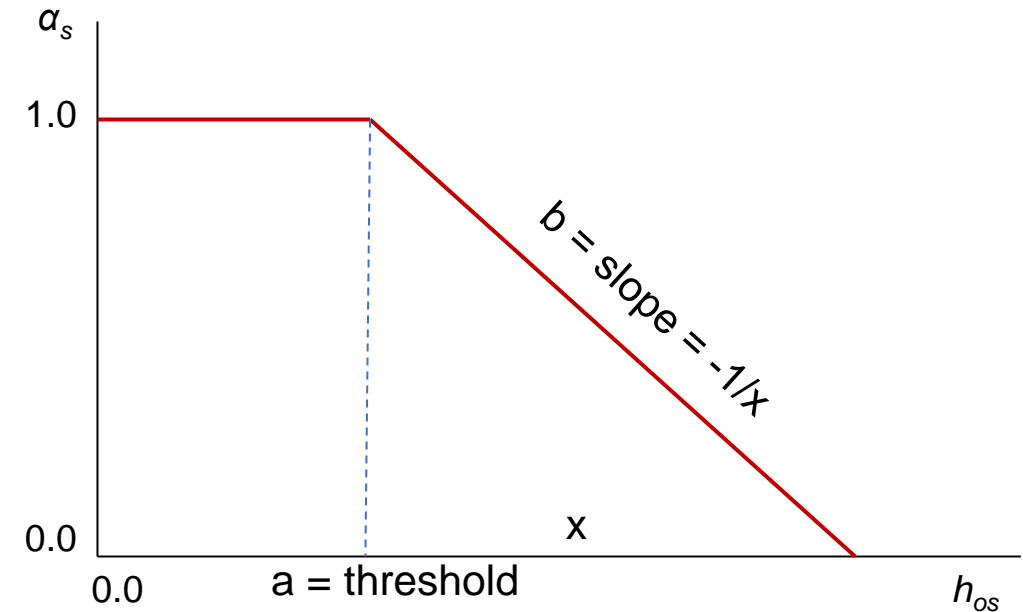
$$\alpha(h) = \frac{1}{1 + \left(\frac{h}{h_{50}}\right)^{p_1}}$$

van Genuchten Uptake reduction factor



## Salinity stress Uptake reduction function

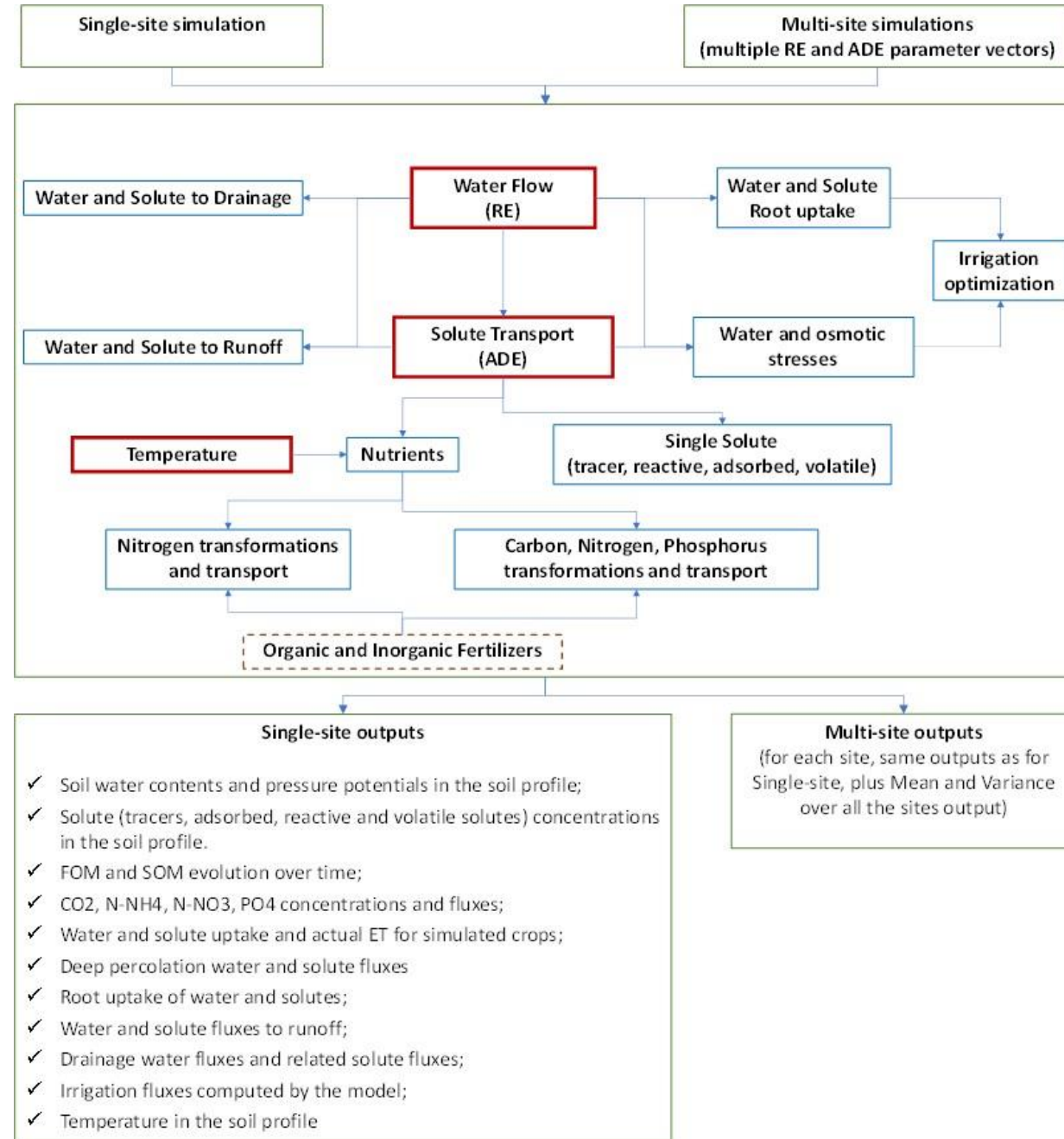
$$\alpha_s(h_{os}) = \begin{cases} 1, & a \leq h_{os} \leq 0 \\ 1 + b(h_{os} - a), & a > h_{os} > a - \frac{1}{b} \\ 0, & h_{os} \leq a - \frac{1}{b} \end{cases}$$



## Multiplicative water and osmotic stress Uptake reduction function

$$\alpha(h, h_{os}) = \frac{1}{1 + \left(\frac{h}{h_{50}}\right)^{p_1}} \frac{1}{1 + \left(\frac{h_{os}}{h_{os,50}}\right)^{p_2}}$$

# Flow Chart of the Processes Calculations



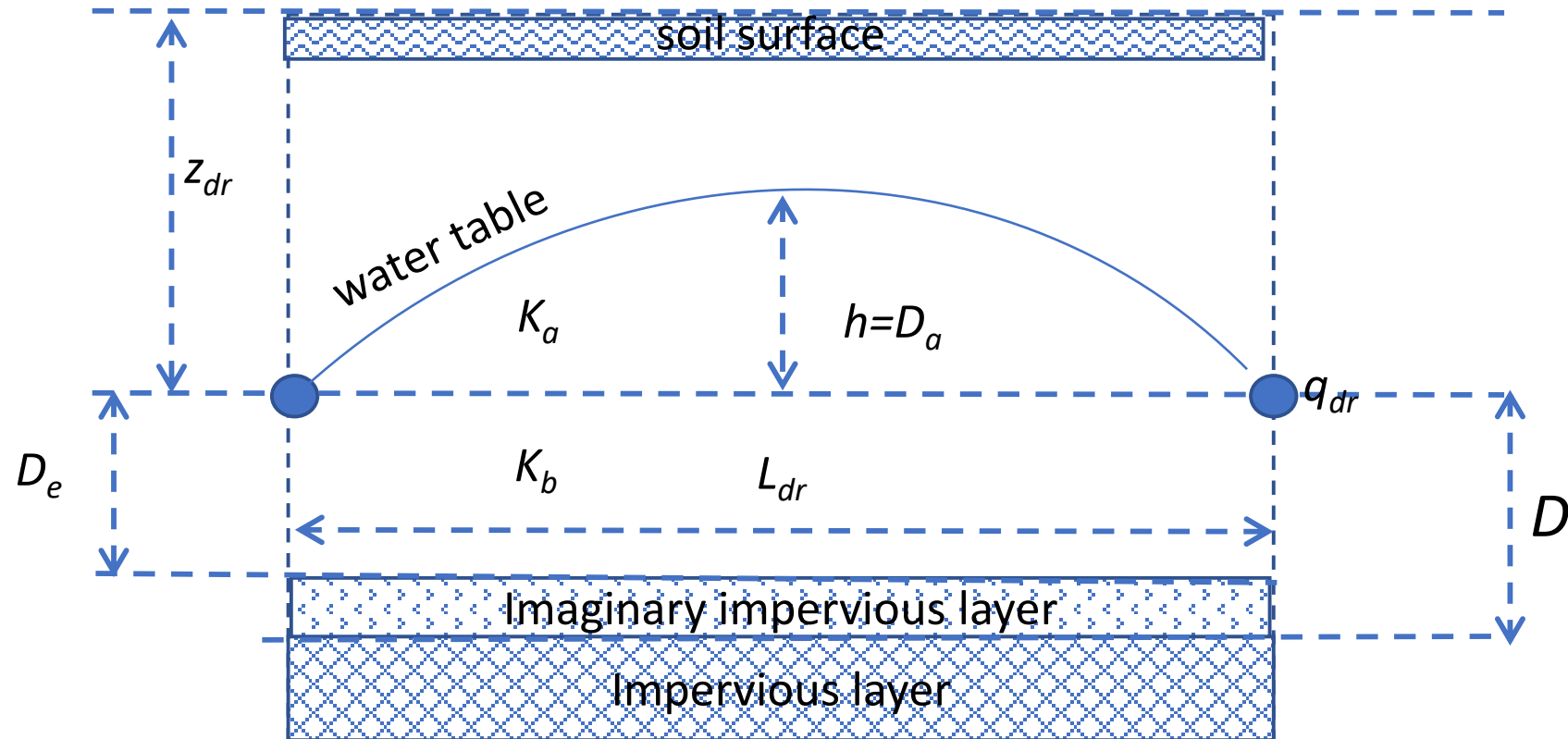
# HOOGHOUDT'S DRAINAGE EQUATION

ADJUSTED FOR ENTRANCE  
RESISTANCE

$$q_{dr} = \frac{8K_b D_{eq} D_a + 4K_a D_a^2}{L_{dr}^2}$$

$S_{dr}$  = fraction of  $q_{dr}$  in  
the  $i^{th}$  node

$Ss_{dr} = S_{dr} \cdot C =$  solute  
flux to the drainage



# Irrigation computed by the model - Irrigation criterion

